



GRUPPO NAZIONALE DI GEOFISICA DELLA TERRA SOLIDA
33° Convegno Nazionale Bologna

Numerical modelling of Self-Potential for Enhanced Geothermal System

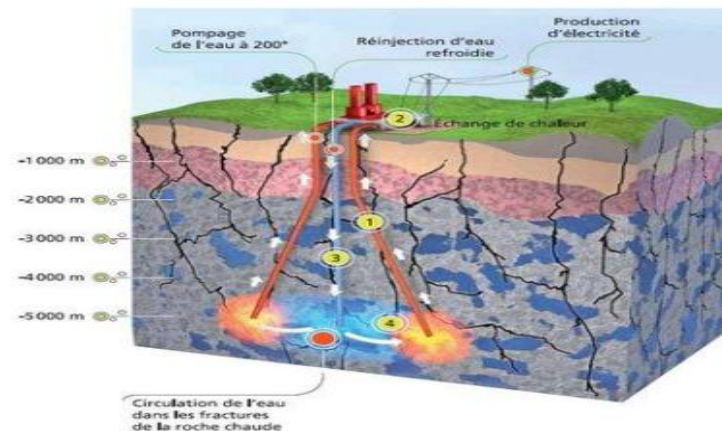
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Introduction

- ▶ Importance of geothermal resource : low environmental impact– reasonable investment
- ▶ EGS system (Enhanced Geothermal System)

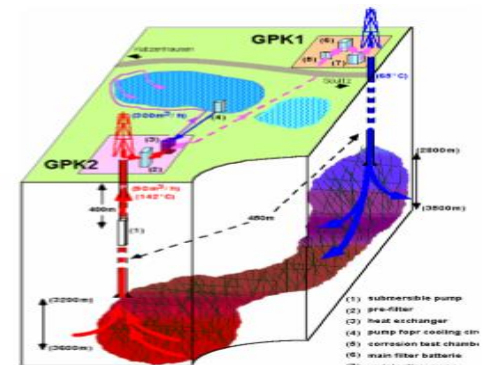
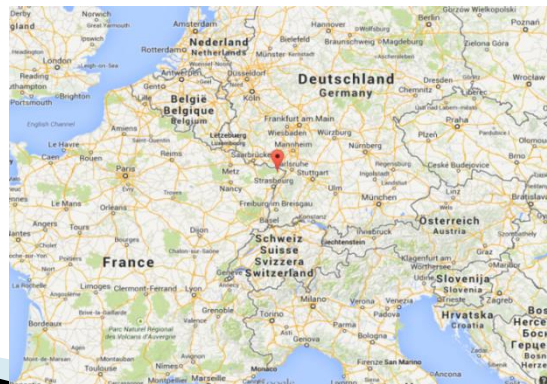


Introduction

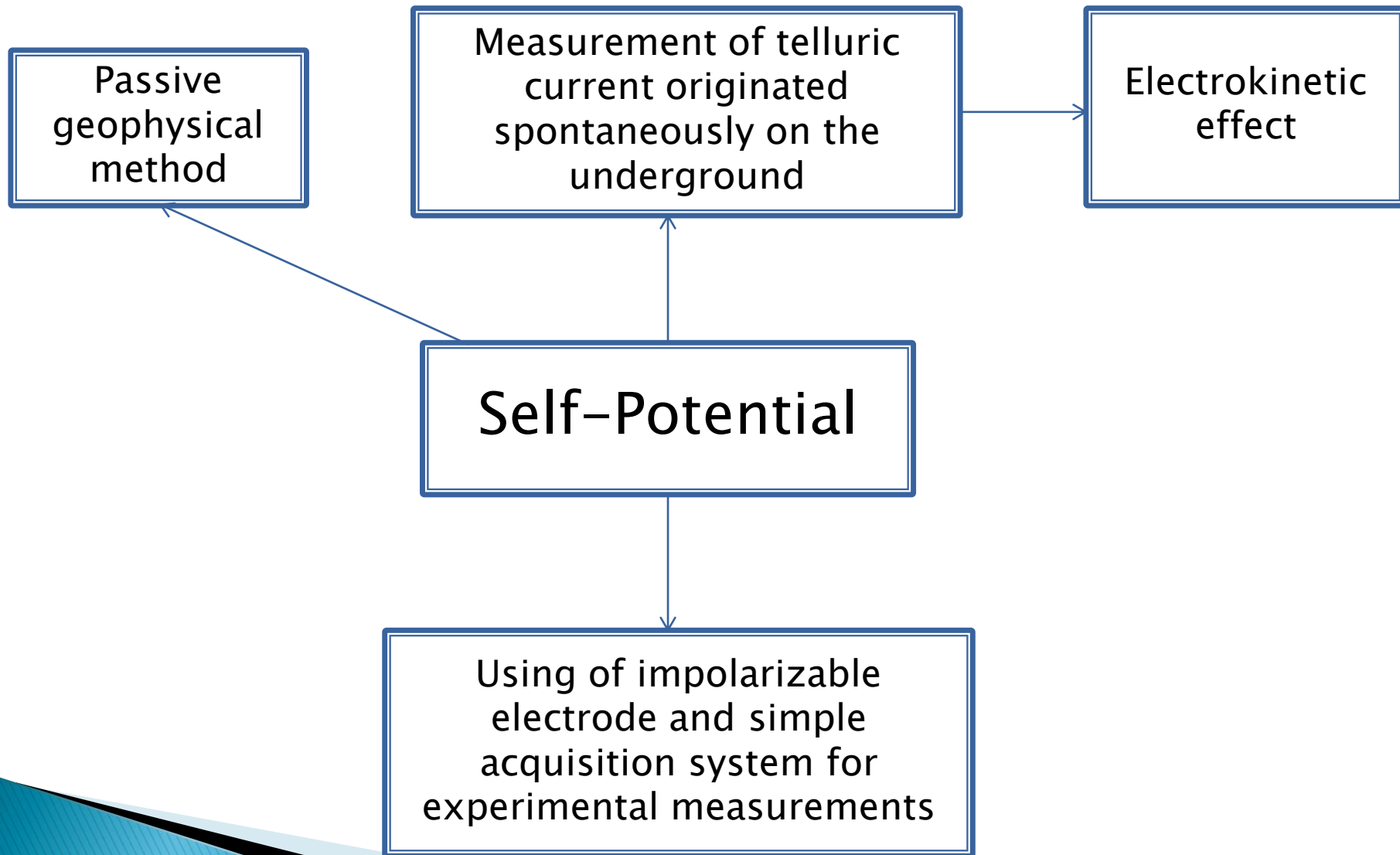
- ▶ EGS exploitation is still perceived as environmentally threatening, because of the problems posed by unwanted induced seismicity above a certain magnitude threshold (MIT Report, 2006)
- ▶ Application of Self-Potential

Introduction

- ▶ Synthetic case: Application of Self-Potential to Soultz-sous-Forets considering one well stimulation to verify the reliability of the model created with Comsol.
- ▶ Real case: Application of Self-Potential to Geothermal area of Soultz-sous-Forets considering a real pumping cycle.

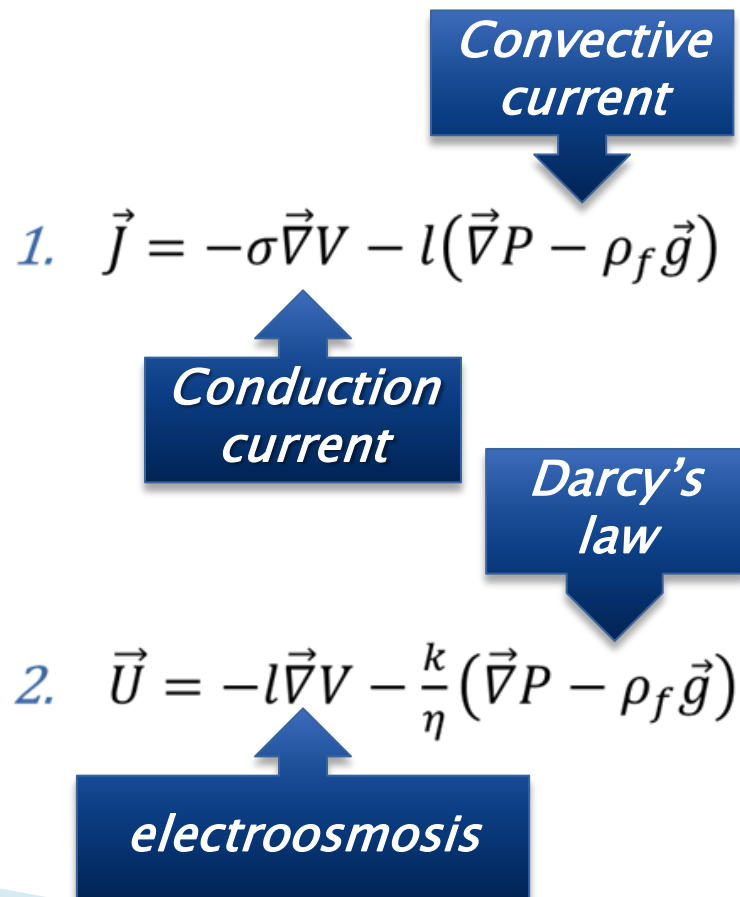


Theory of Self-Potential



Theory of Self-Potential

- ▶ Equation used for our model are:



Numerical procedure

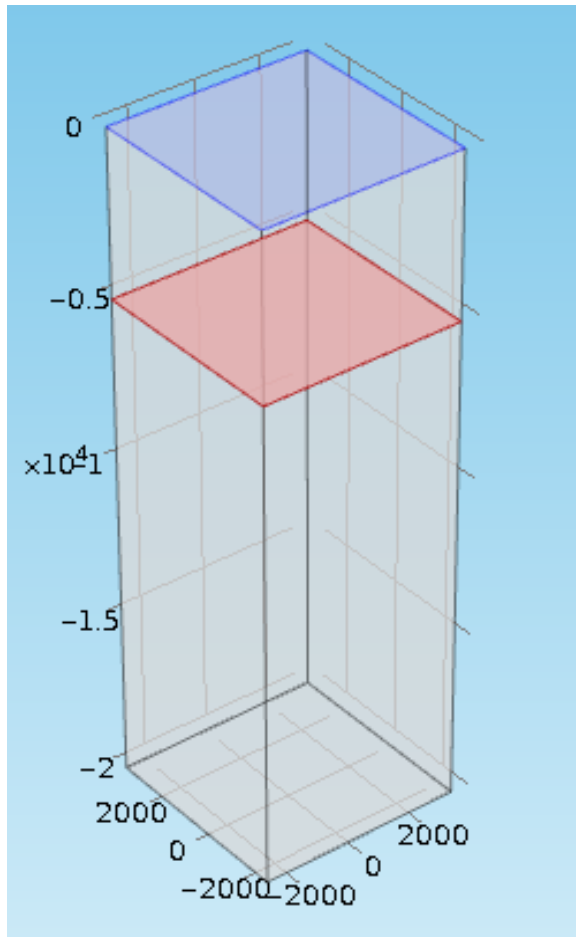
- ▶
- ▶ The variation of pressure retrieved with Tough 2 has been imposed as source in a Self-Potential model created in Comsol for both the case under study
- ▶ The model resolves the Poisson's equation:

$$\nabla^2 V = -\frac{\vec{\nabla} \sigma}{\sigma} \vec{E} - \frac{1}{\sigma} [\vec{\nabla} l \cdot \vec{\nabla} P - \rho_f \vec{\nabla} l \cdot \vec{g} - l \nabla^2 P]$$

First case: synthetic simulation

- ▶ Achievement of a simulation considering the pumping cycle in one well at fixed injection rate (30 kg/s) for 16 days
- ▶ Goal of simulation:
 - Determination of electric potential at earth's surface
 - Evaluation about the reliability of the model

First case: inside the model



Eq. resolved:

Source :

homogeneous conductivity model ($\sigma = 0.001 \text{ S/m}$)

$= C \cdot \sigma$ (electrokinetic coupling coefficient)

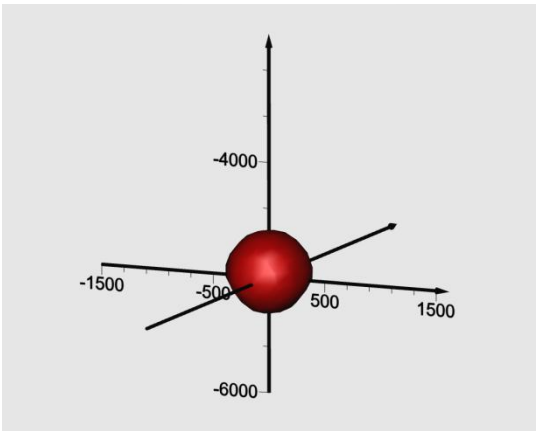
$C = 2 \text{ mV/bar}$ (Révil et al. 1999)

First case: Isosurface of pressure

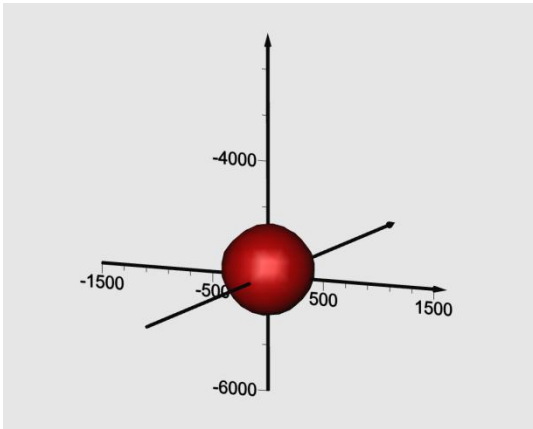


DP=10 bar

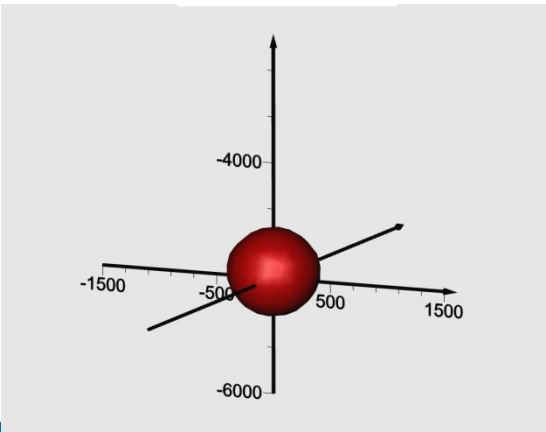
Day 20



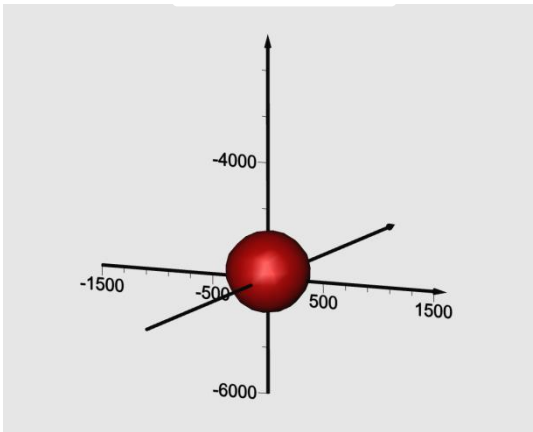
Day 30



Day 50



Day 60

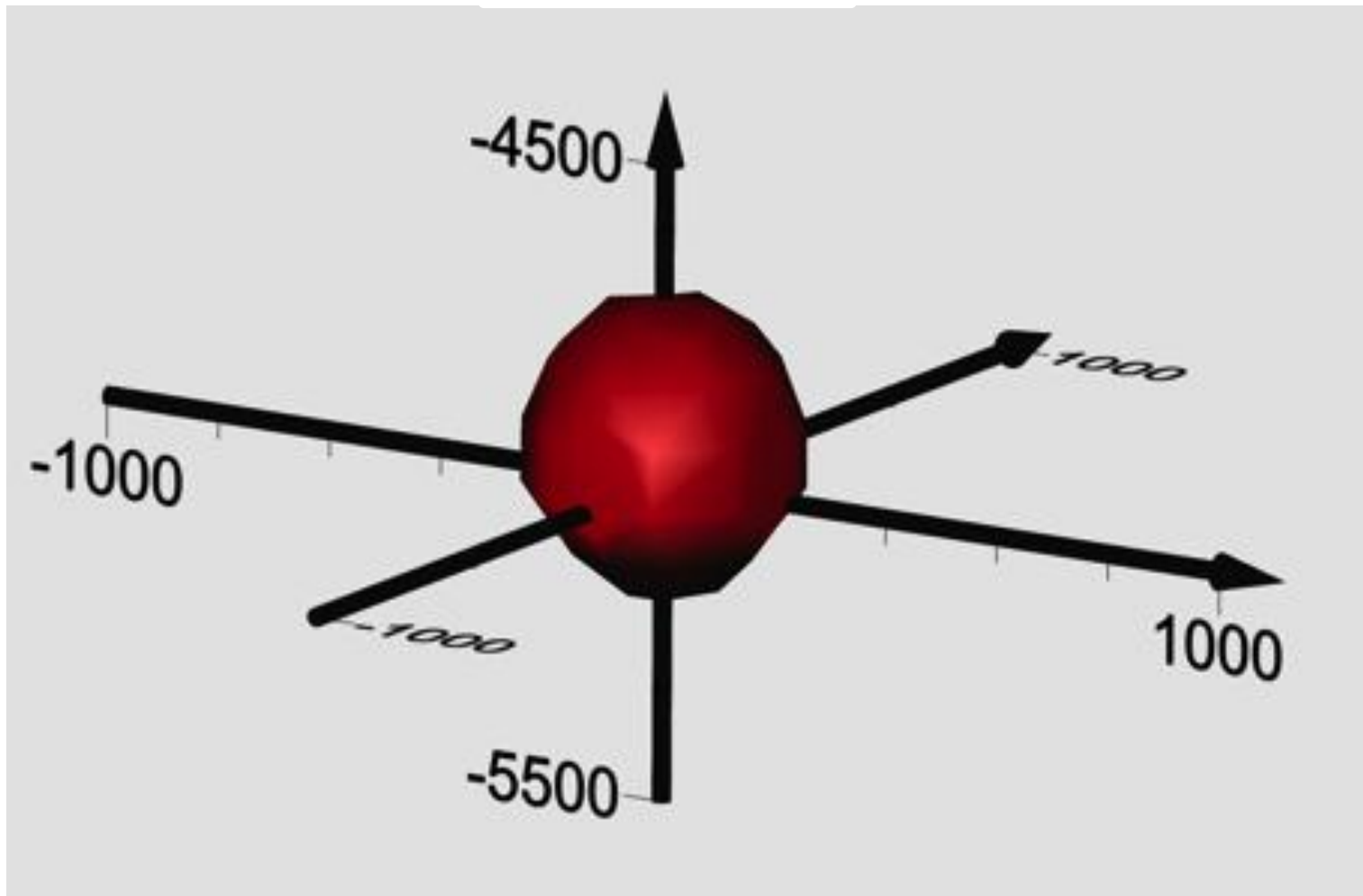


First case: Isosurface of pressure




DP=10 bar

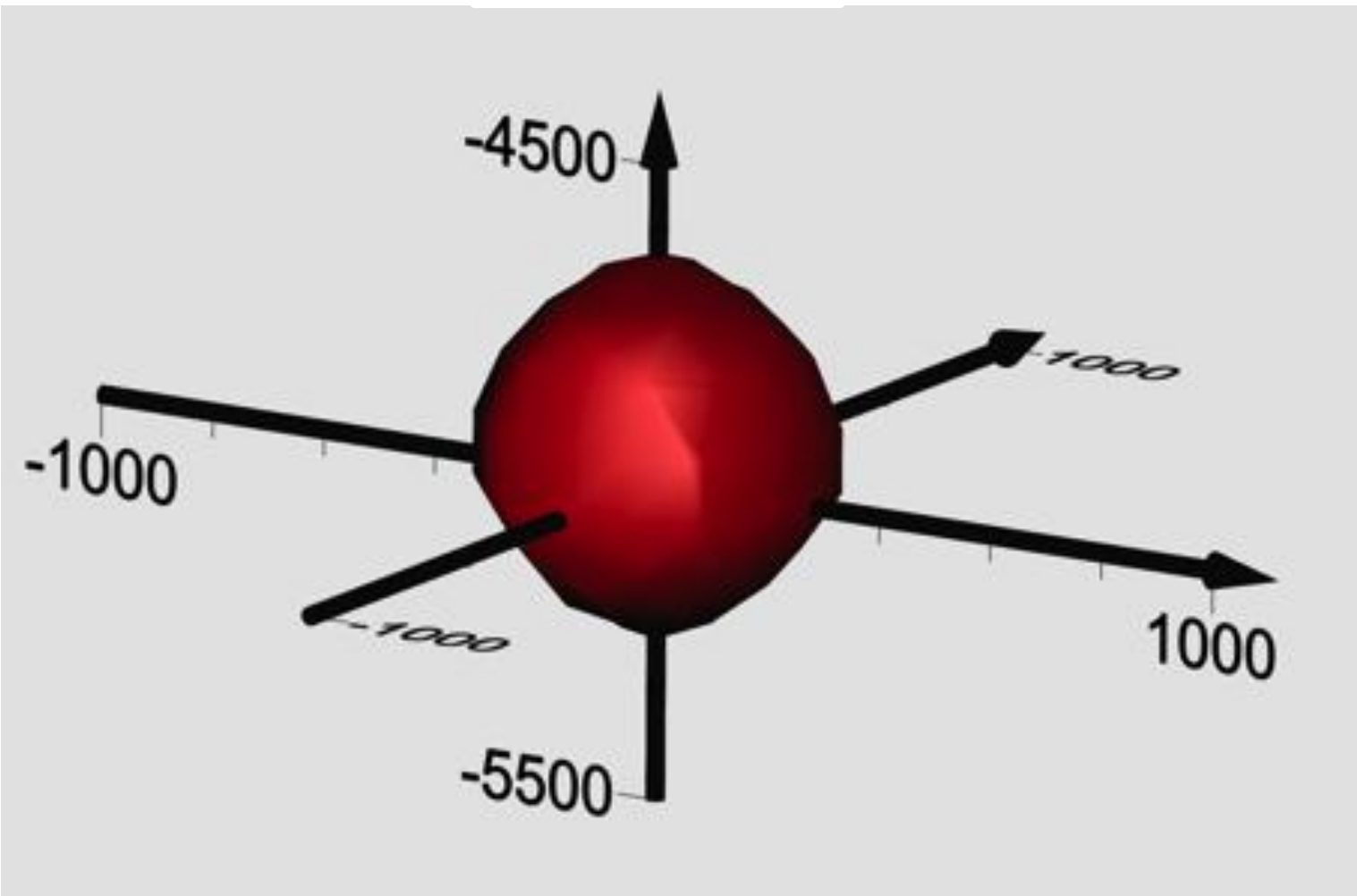
Day 5



First case: Isosurface of pressure

 DP=10 bar

Day 10

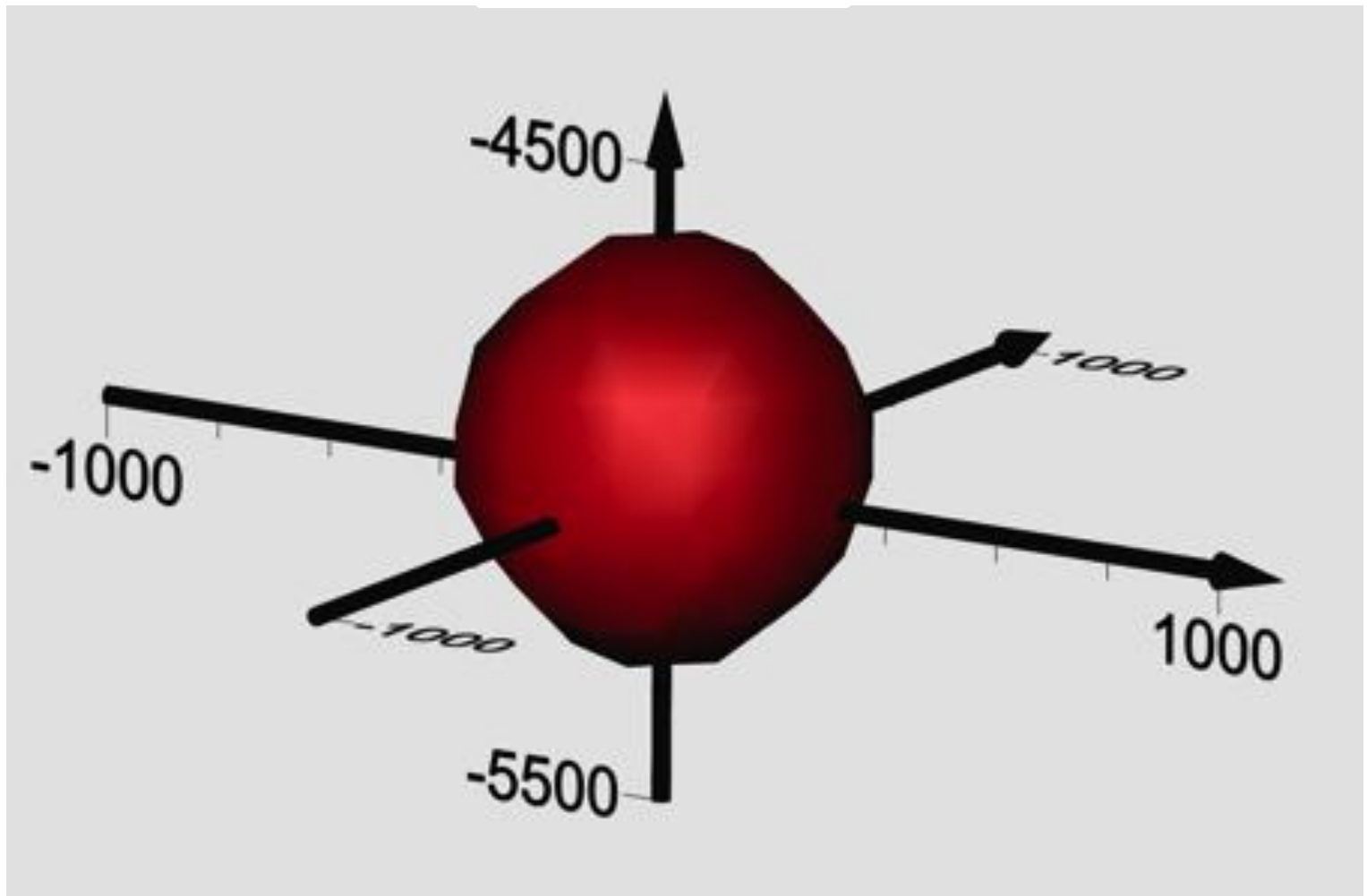


First case: Isosurface of pressure



DP=10 bar

Day 15

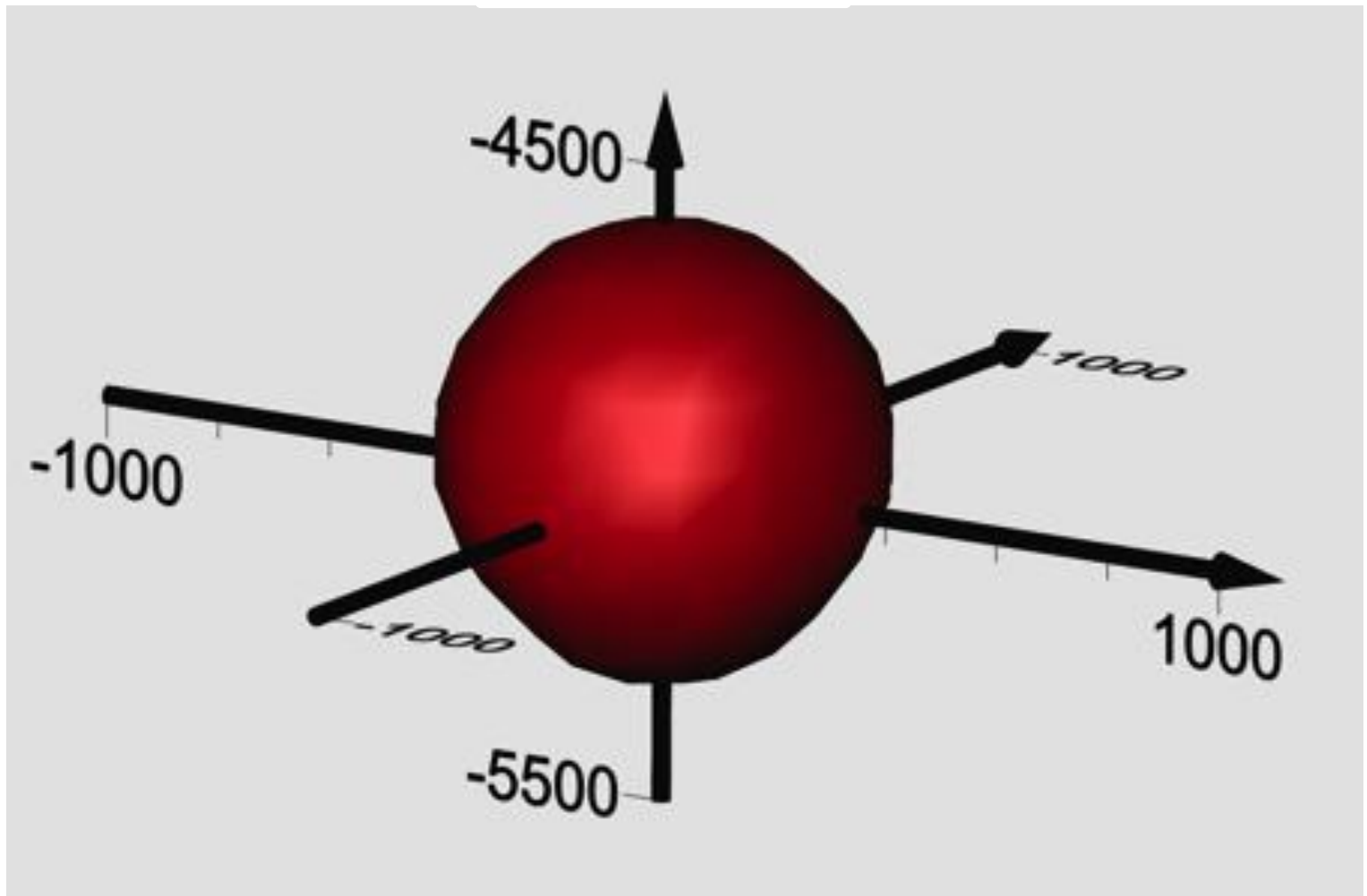


First case: Isosurface of pressure



DP=10 bar

Day 20

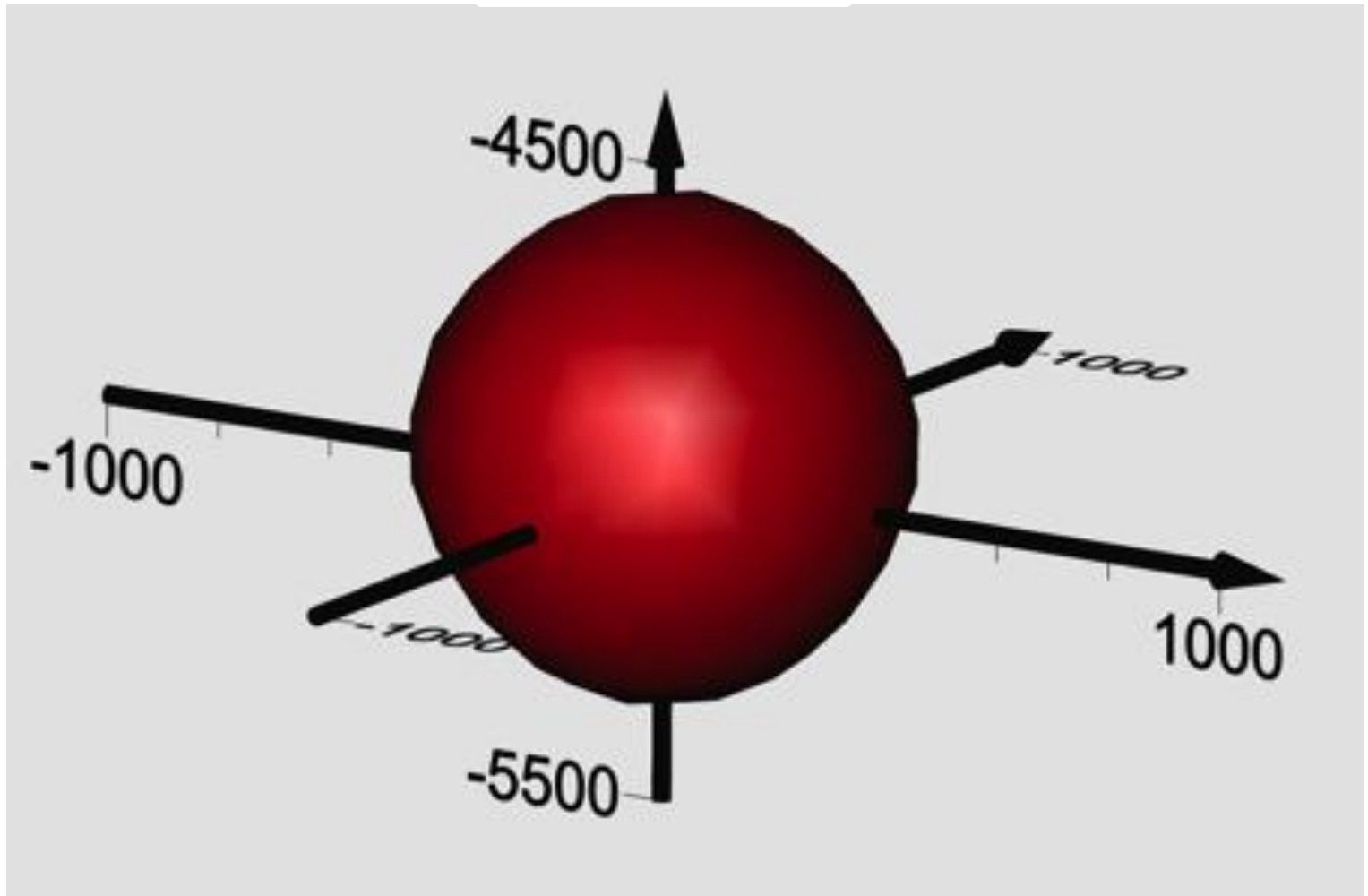


First case: Isosurface of pressure



DP=10 bar

Day 30

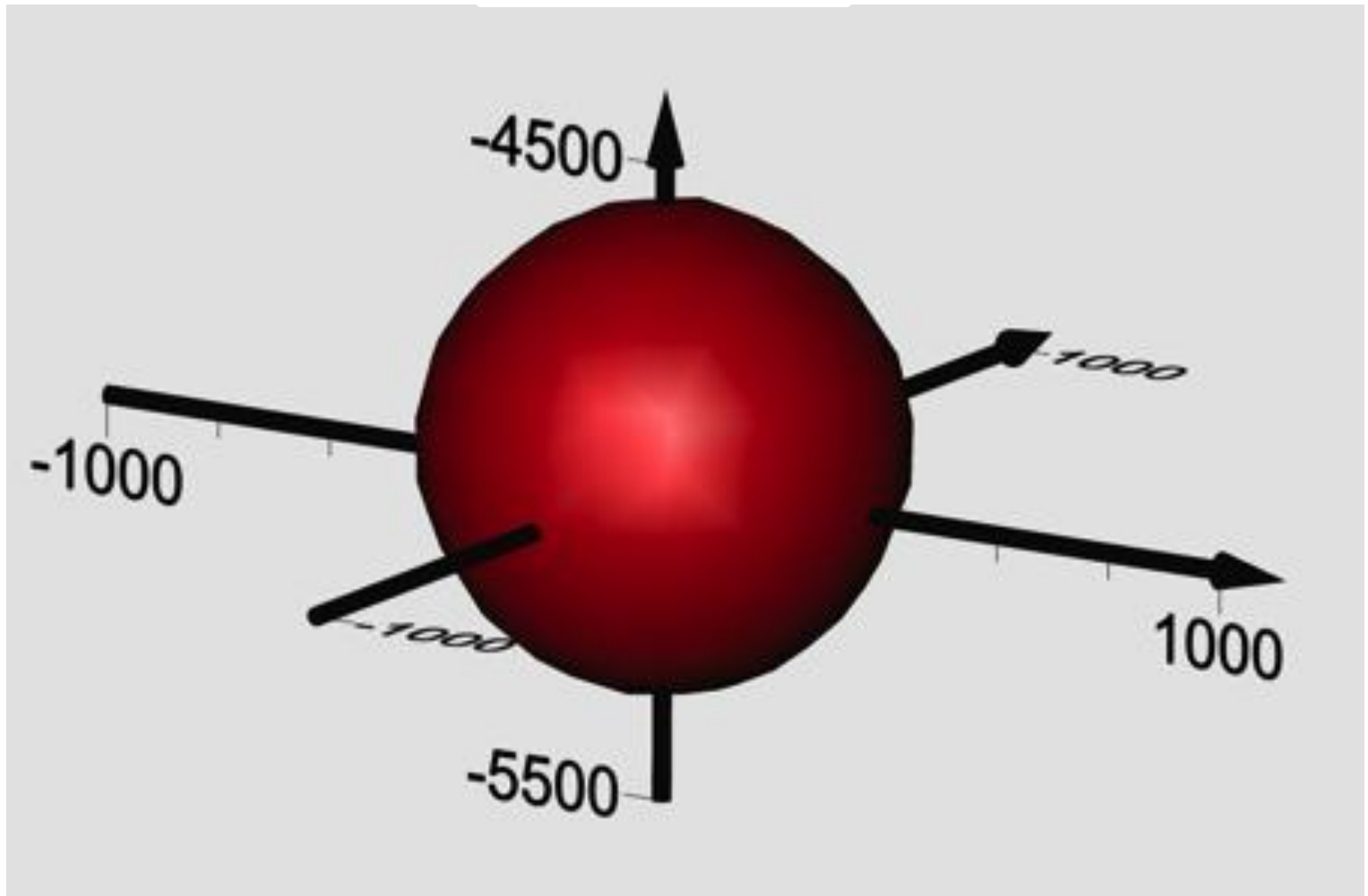


First case: Isosurface of pressure



DP=10 bar

Day 50

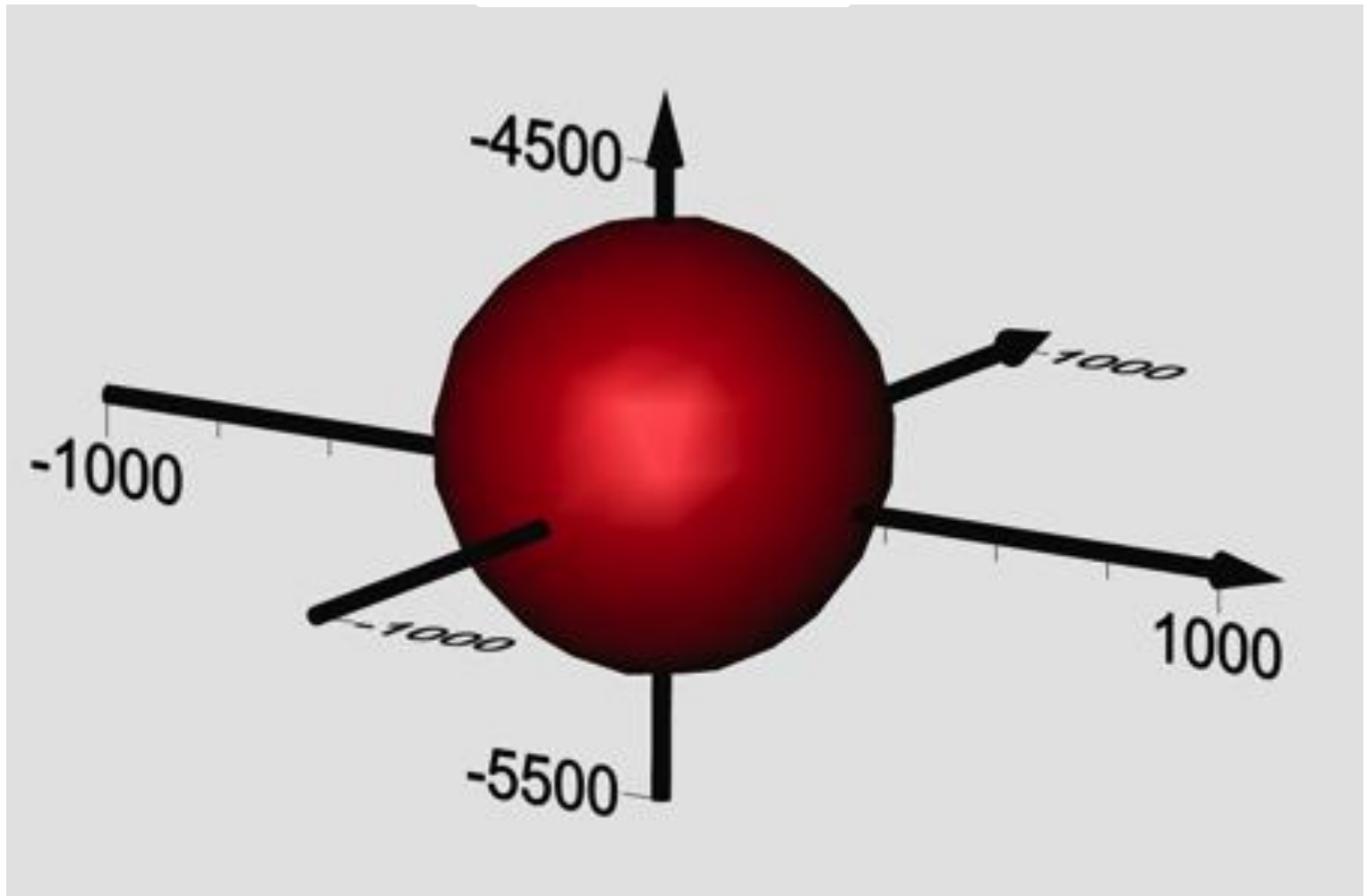


First case: Isosurface of pressure



DP=10 bar

Day 60

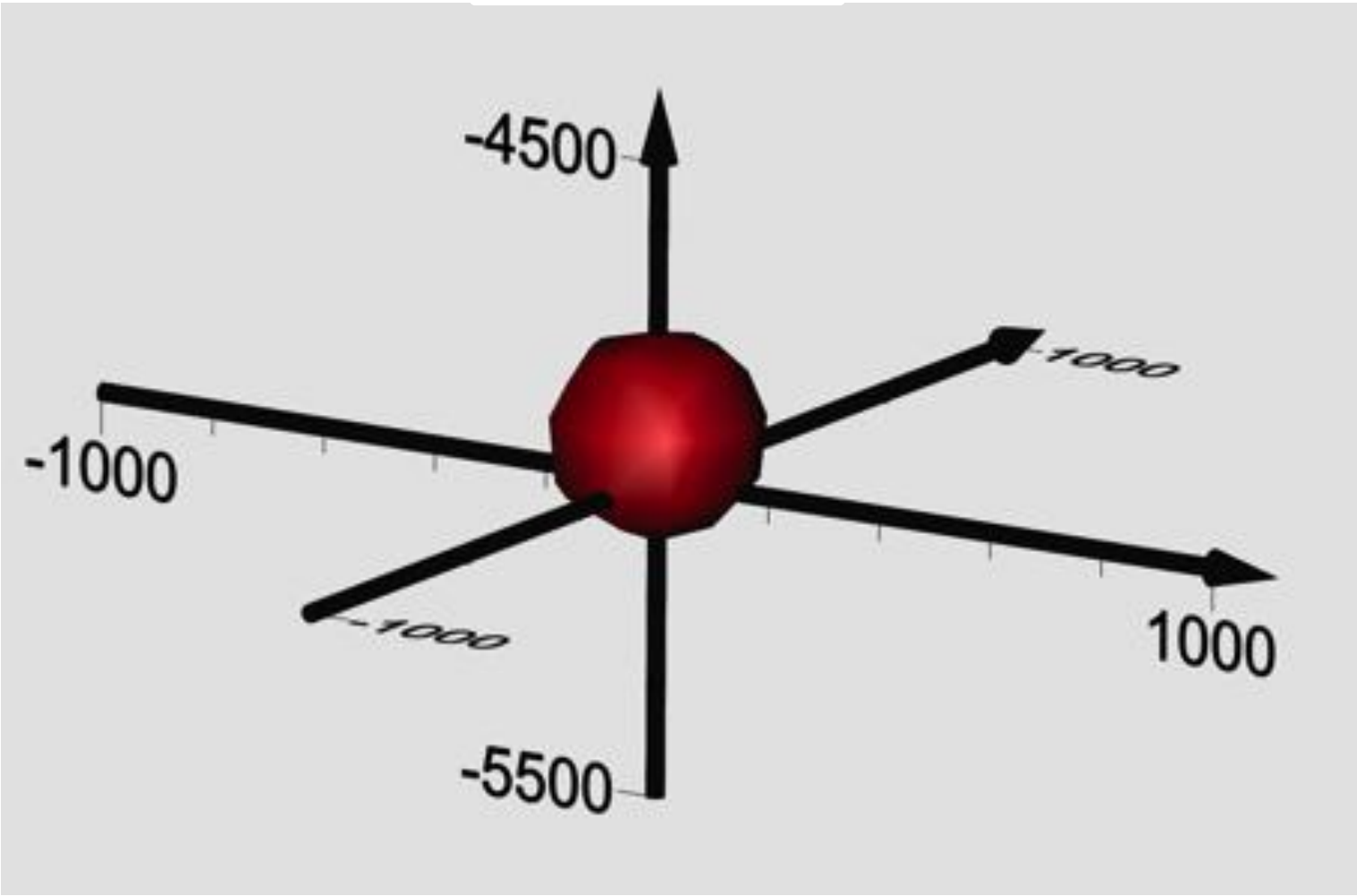


First case: Isosurface of pressure

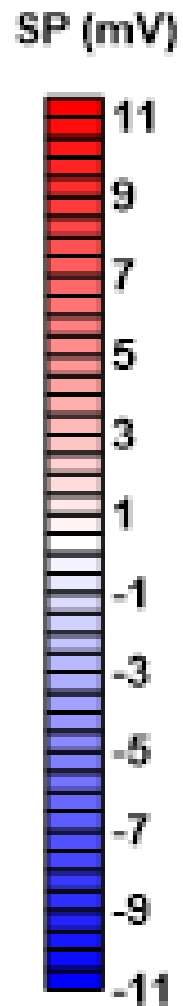
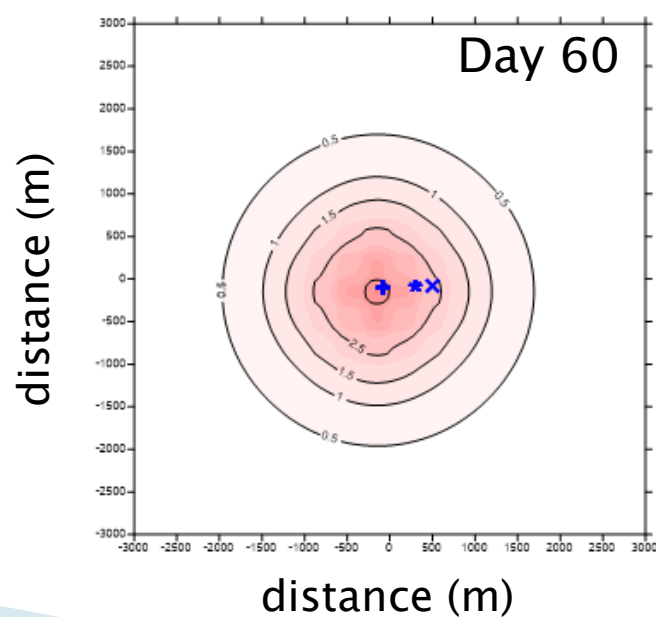
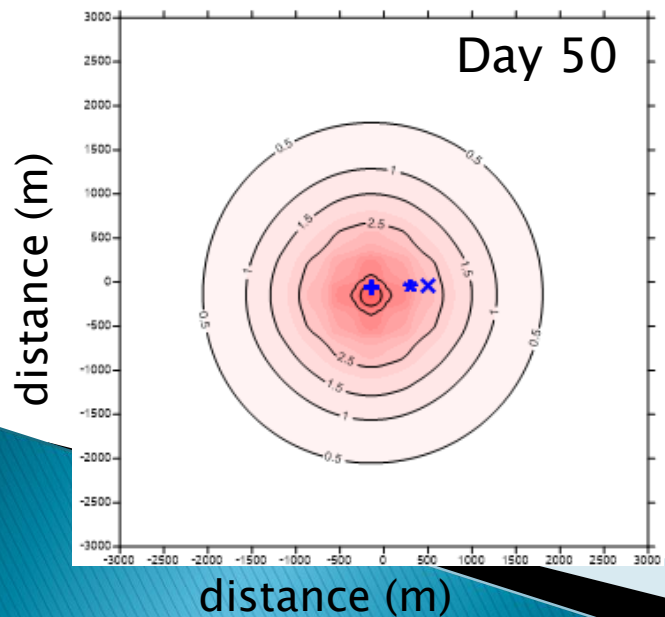
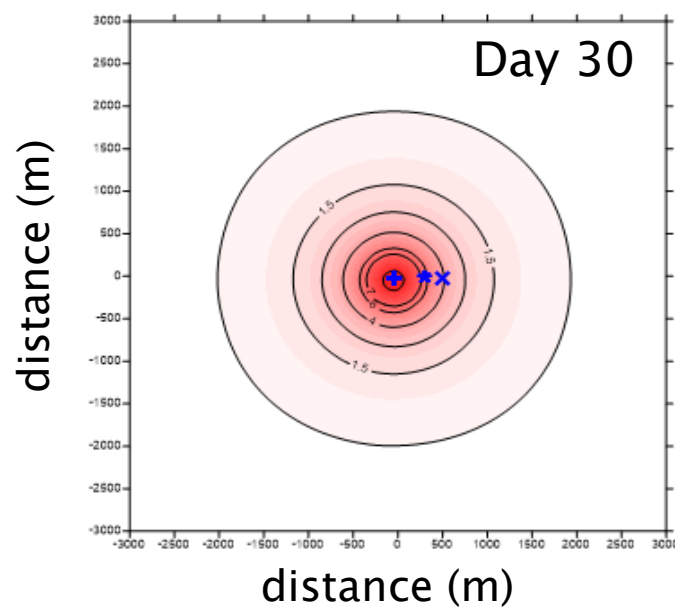
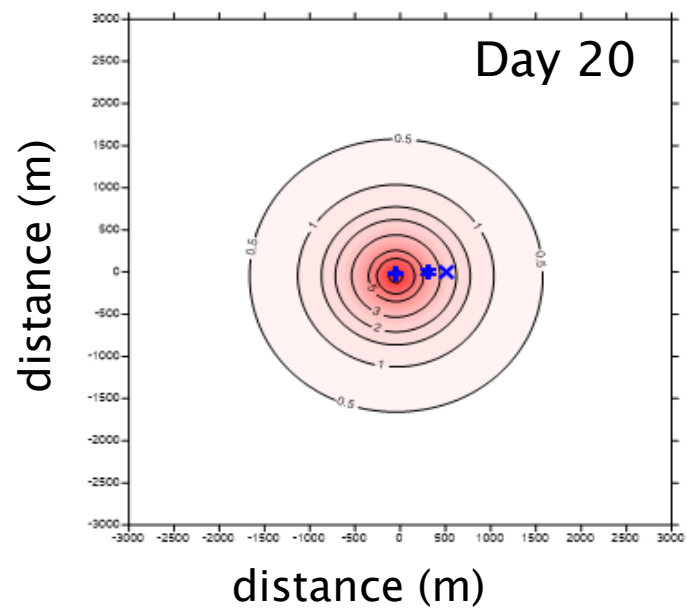


DP=10 bar

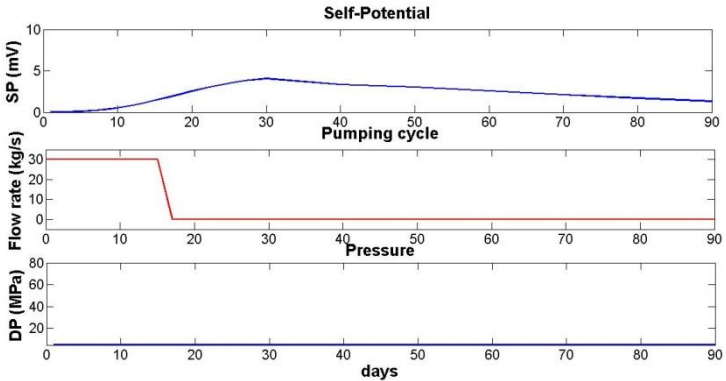
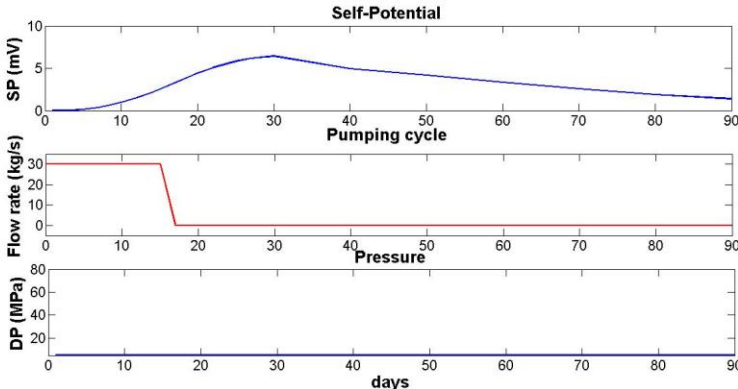
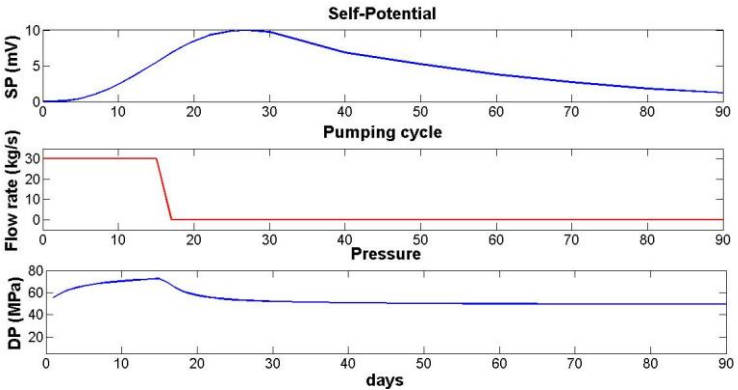
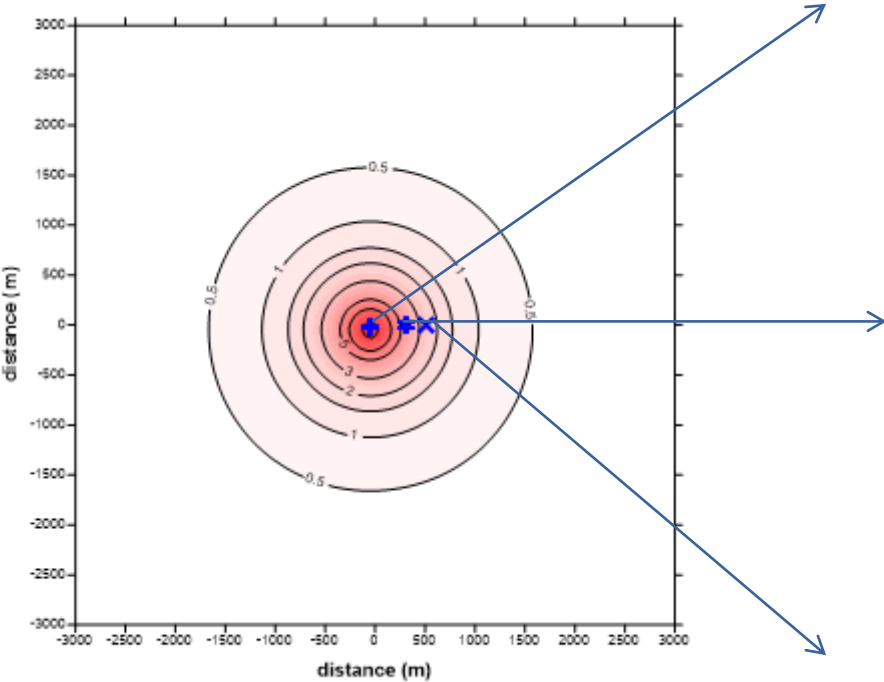
Day 90



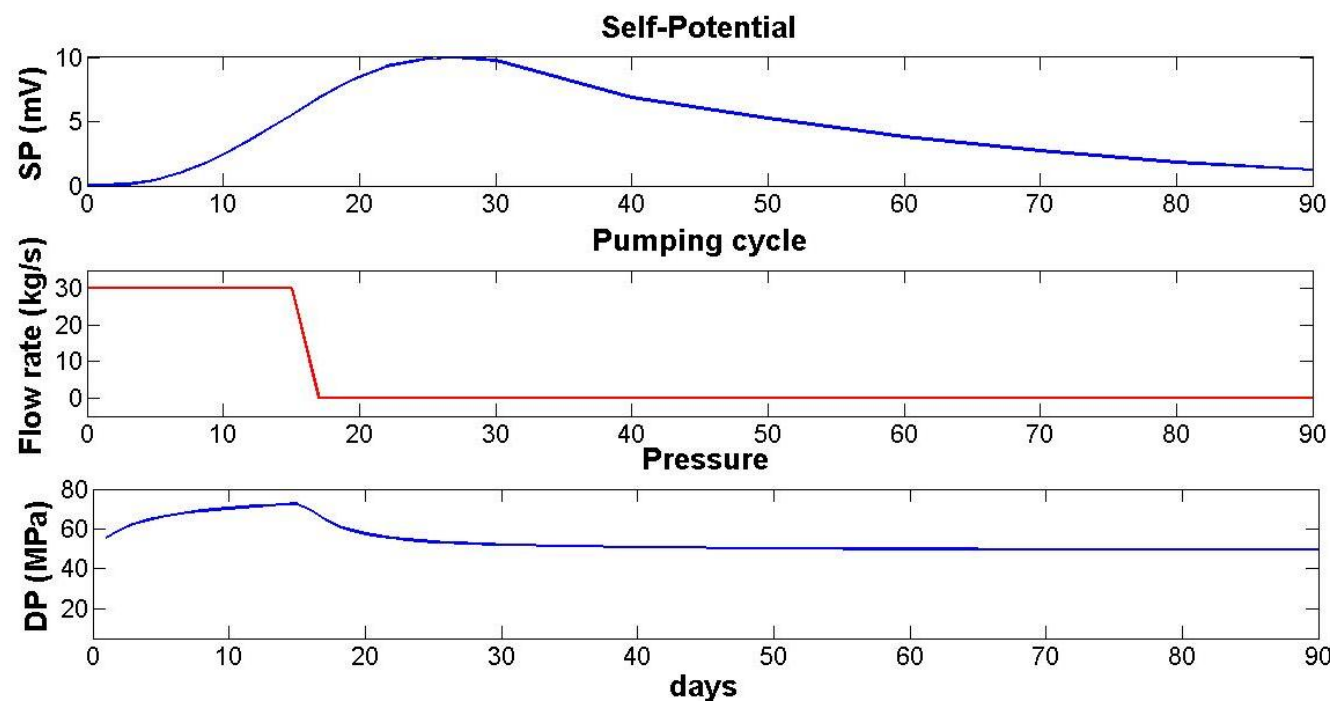
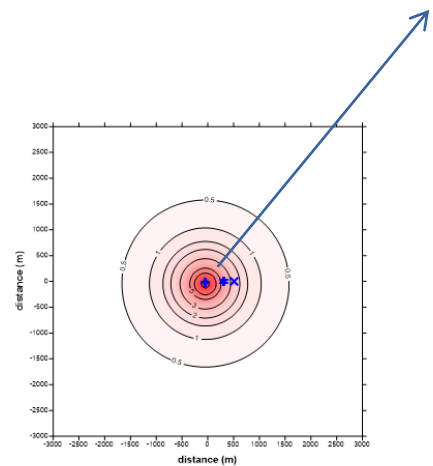
First case: electric potential distribution at earth's surface



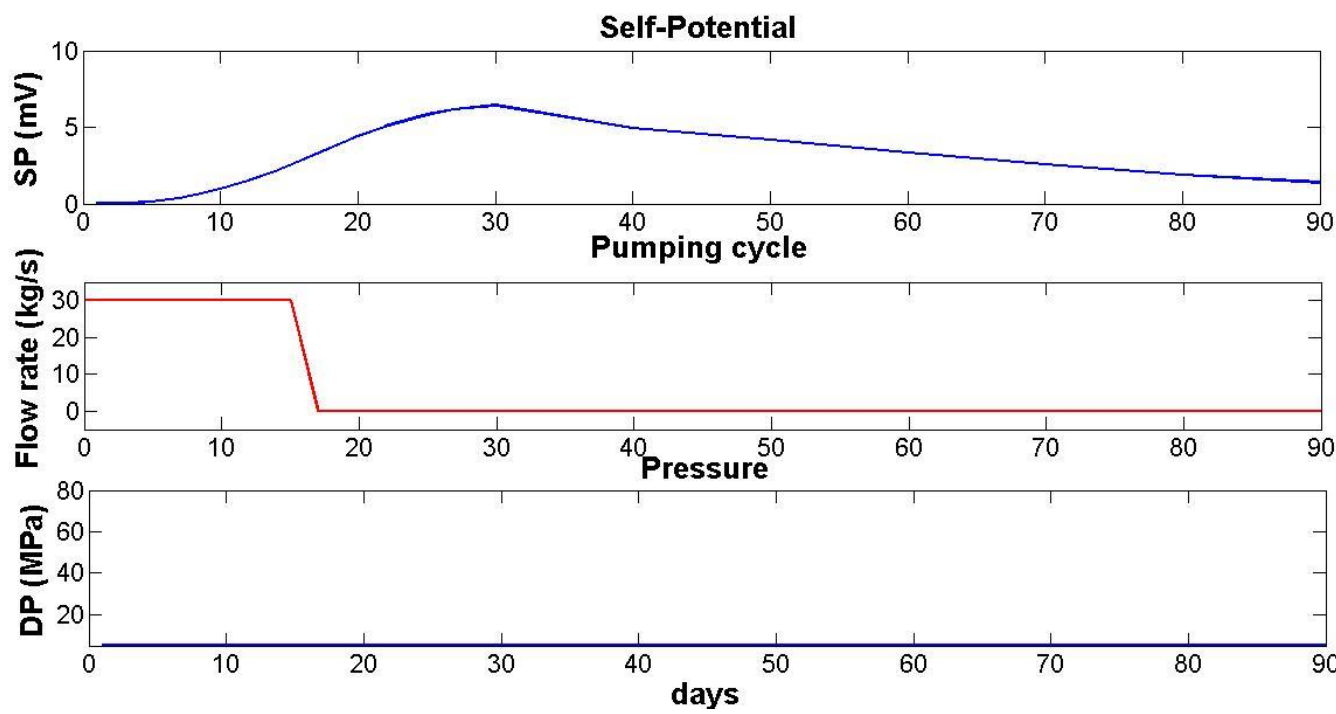
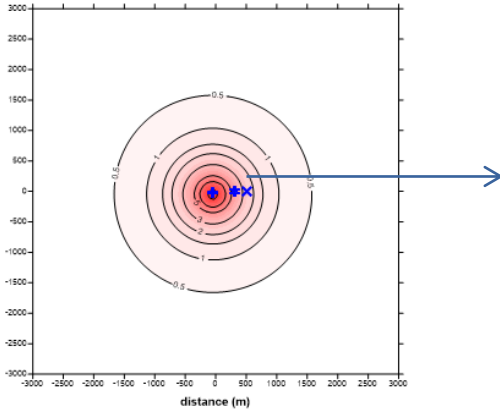
first case: time dependent trend of electric potential



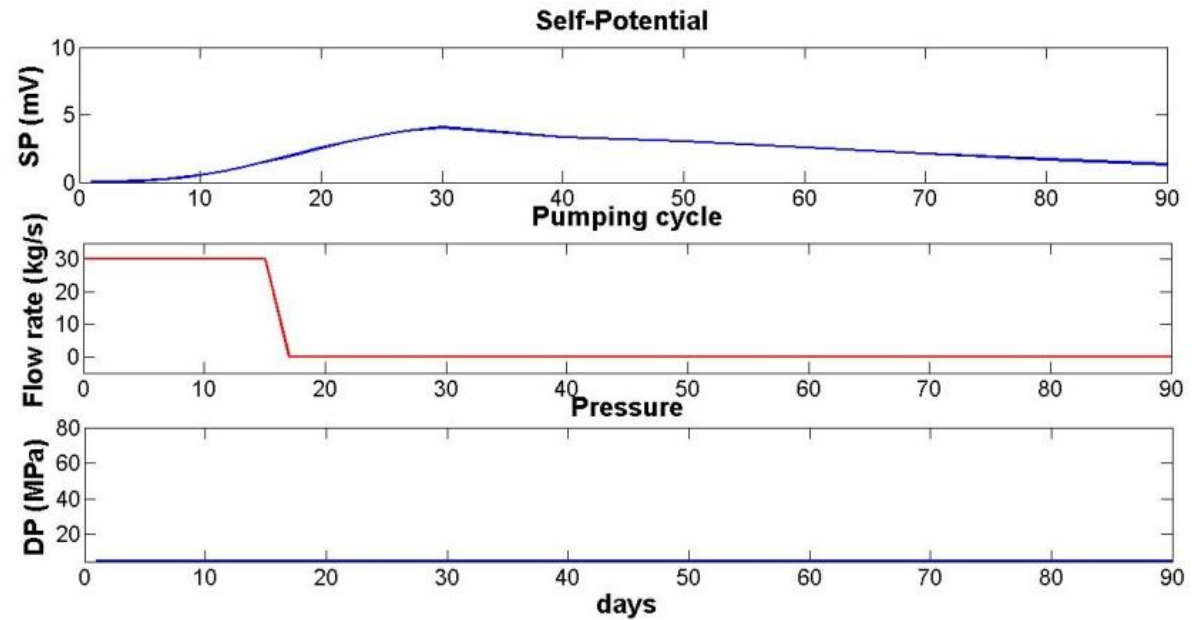
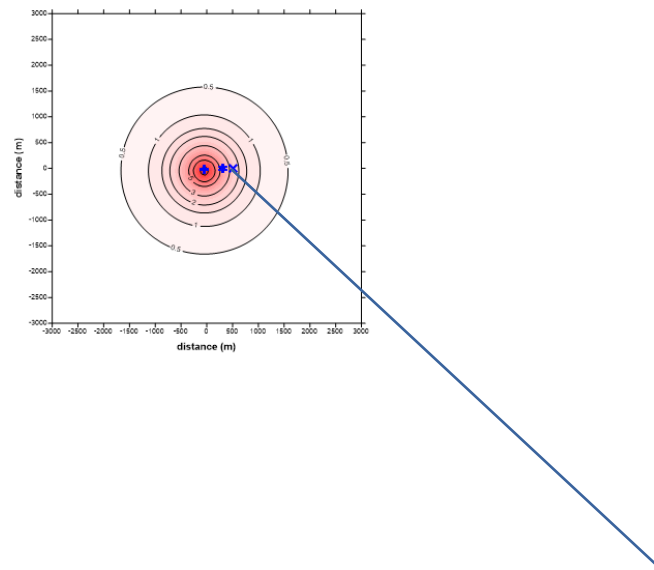
first case: time dependent trend of electric potential



first case: time dependent trend of electric potential



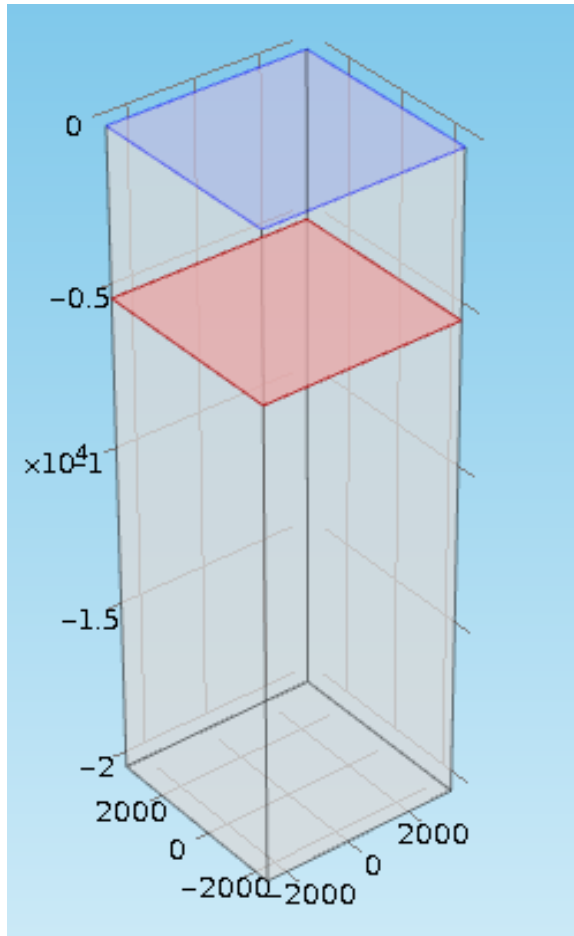
first case: time dependent trend of electric potential



Second case: Real simulation

- ▶ Achievement of a simulation considering the pumping cycle in two wells at different injection rates (30 kg/s–50 kg/s) for 12 days
- ▶ Goal of simulation:
 - Determination of electric potential at earth's surface
 - Underline the difference between the trend of pressure and electric potential
 - Search a connection between our results and the behaviour of induced seismicity

Second case: inside the model



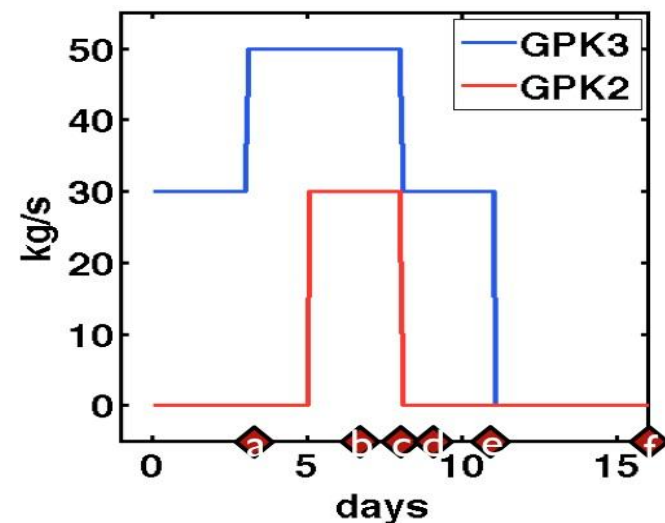
$$\text{Eq. resolved: } \nabla^2 V = -\frac{\vec{\nabla}\sigma}{\sigma} \vec{E} - \frac{1}{\sigma} [\vec{\nabla}l \cdot \vec{\nabla}P - \rho_f \vec{\nabla}l \cdot \vec{g} - l \nabla^2 P]$$

$$\text{Source : } \frac{\vec{\nabla}\sigma}{\sigma} \vec{E} + l \nabla^2 P$$

heterogeneous conductivity model (Geiermann et al.2011)

$l = C * \sigma$ (electrokinetic coupling coefficient)

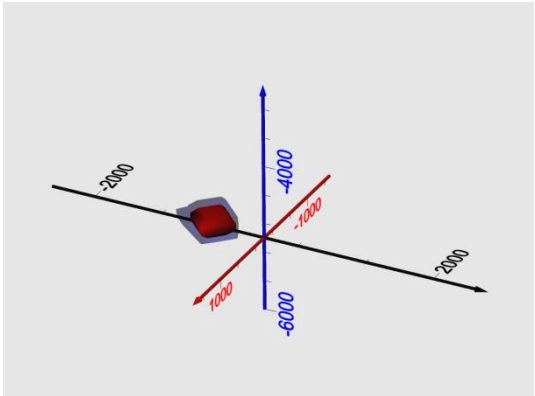
$C = 2 \text{ mV/bar}$ (Révil et al.1999)



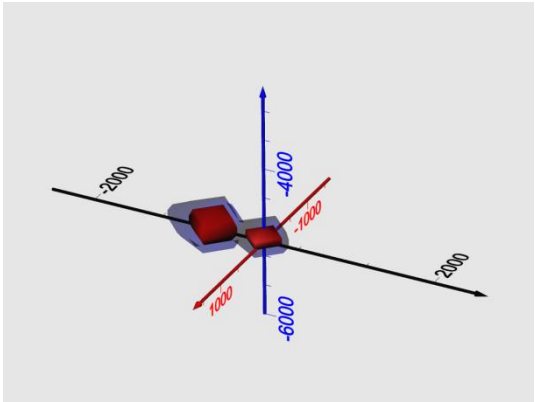
Second case: Isosurface of pressure

DP= 1 bar
DP=10 bar

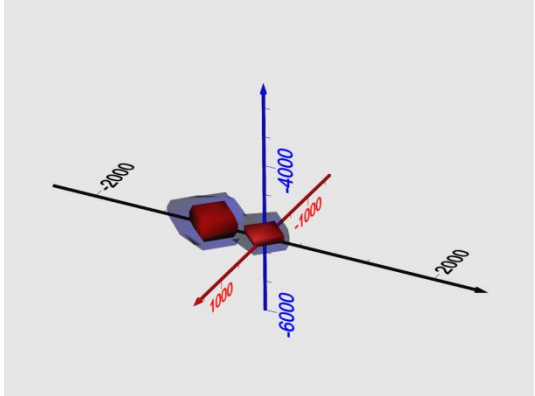
Day 3



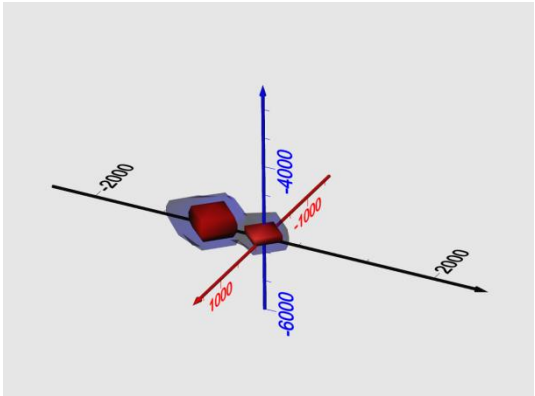
Day 7



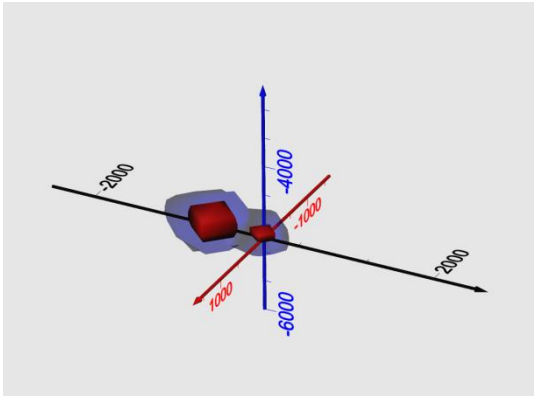
Day 8



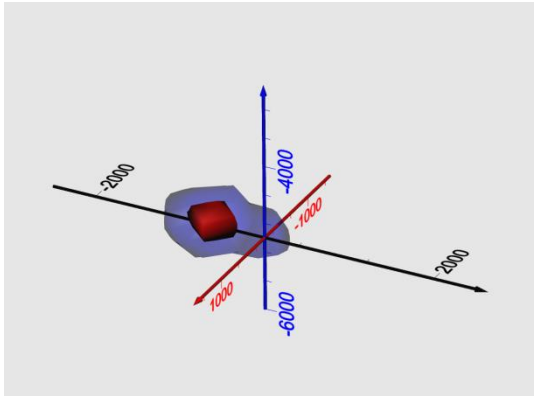
Day 9



Day 12

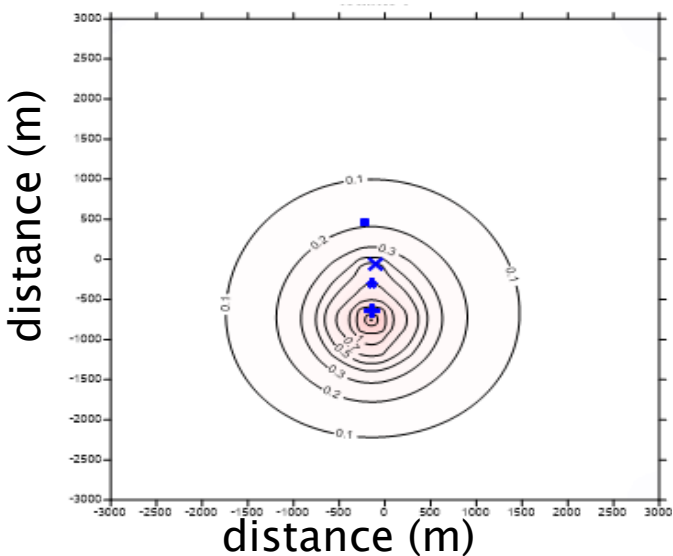


Day 16

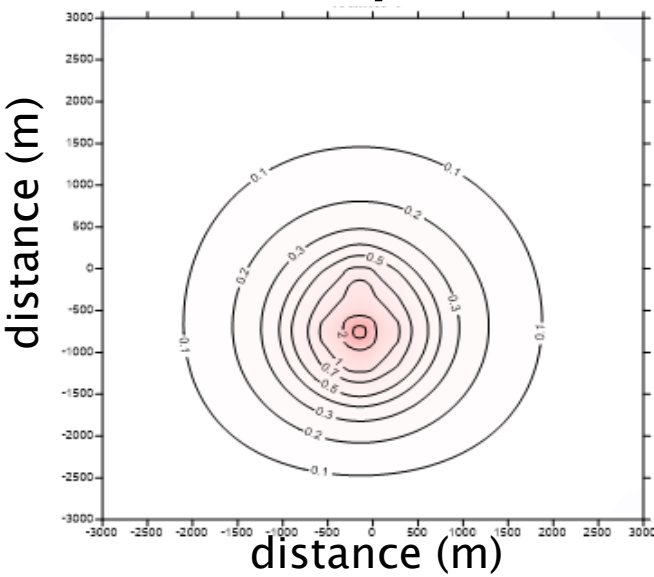


Second case: electric potential distribution at earth's surface

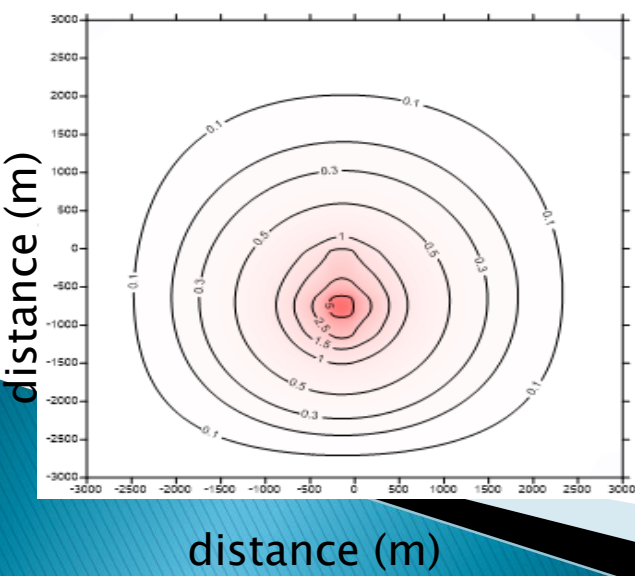
Day 7



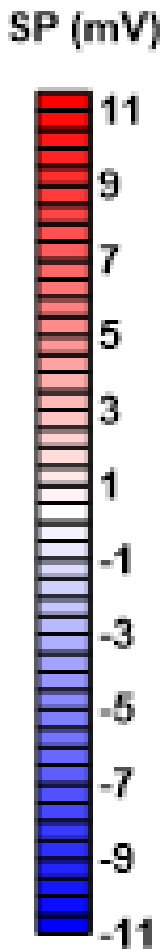
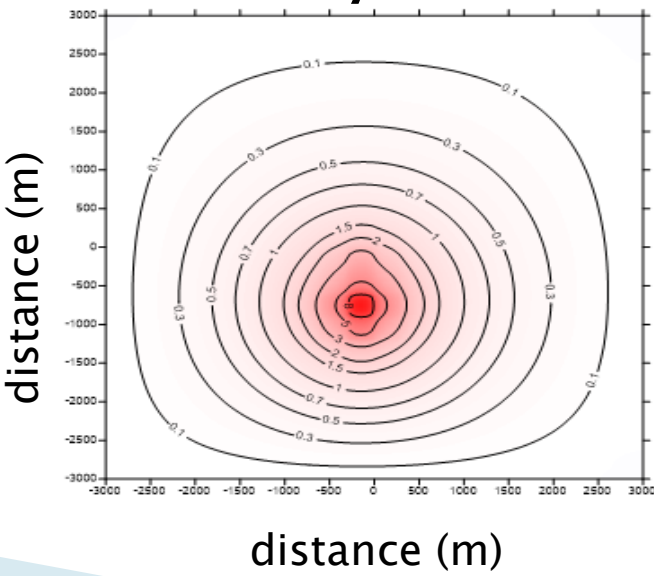
Day 9



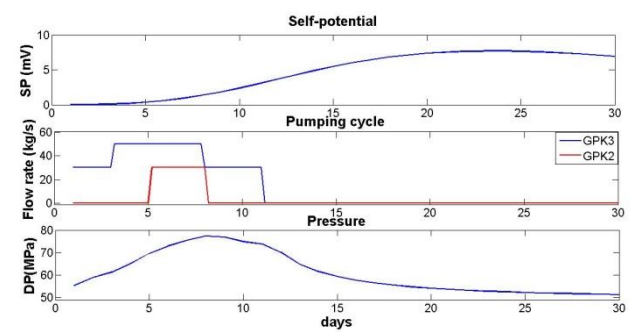
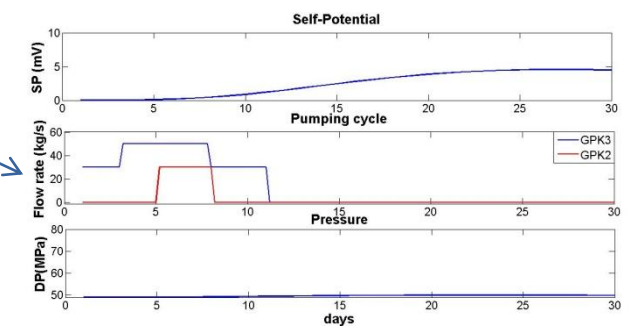
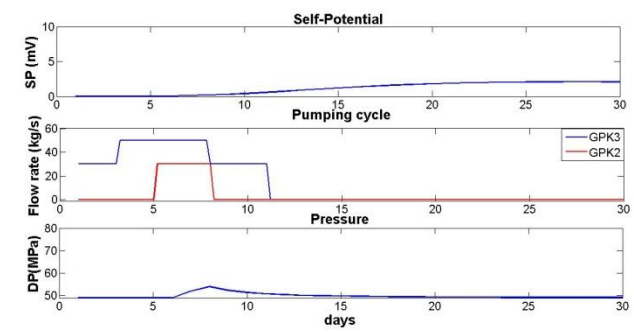
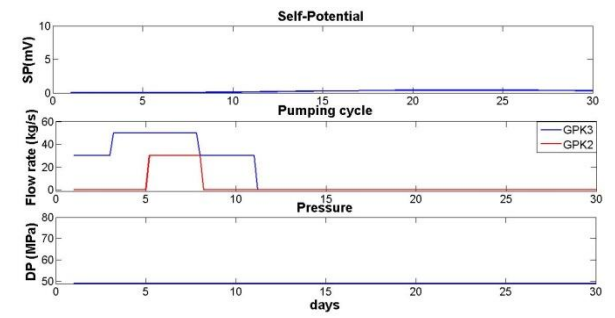
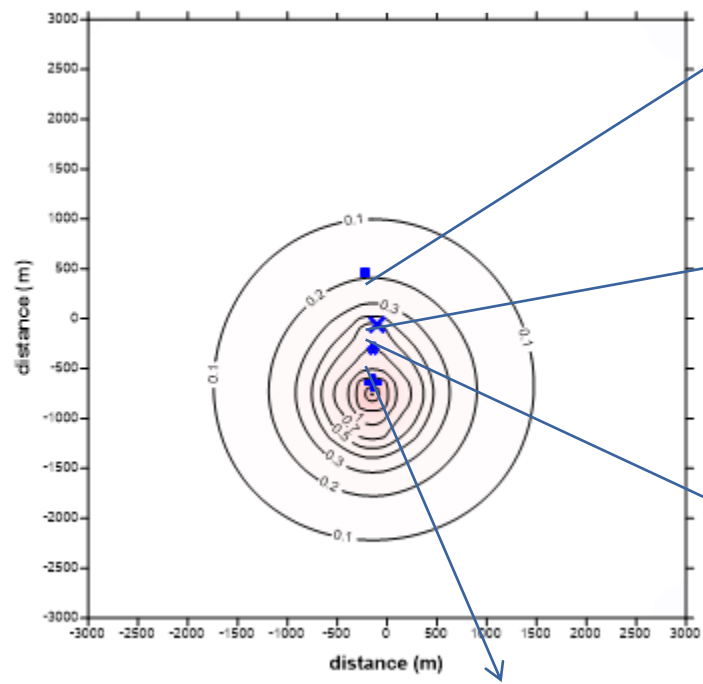
Day 12



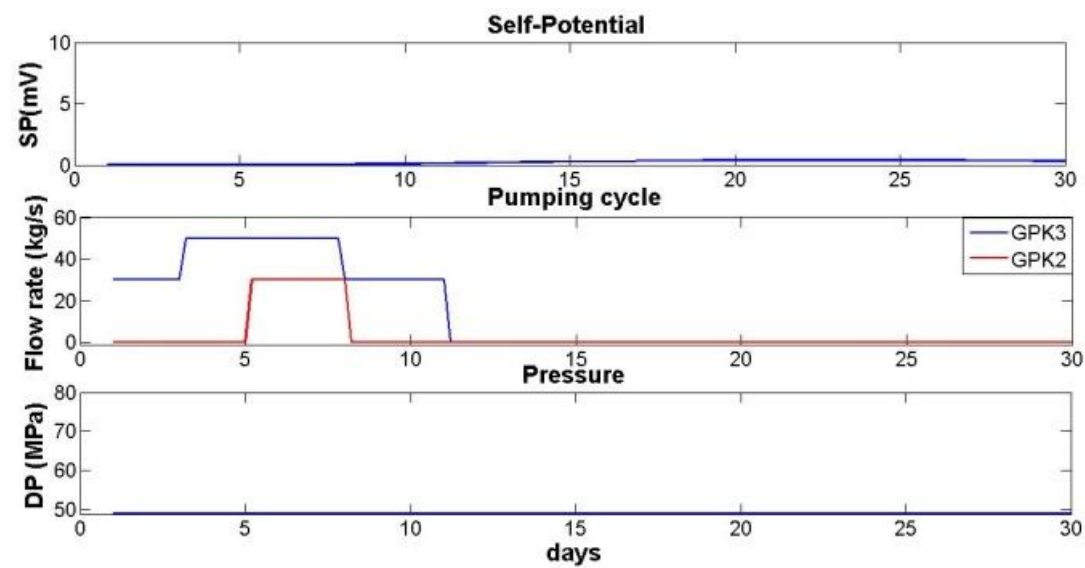
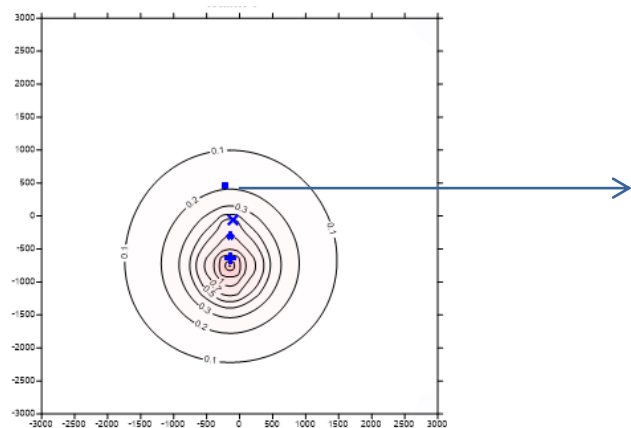
Day 16



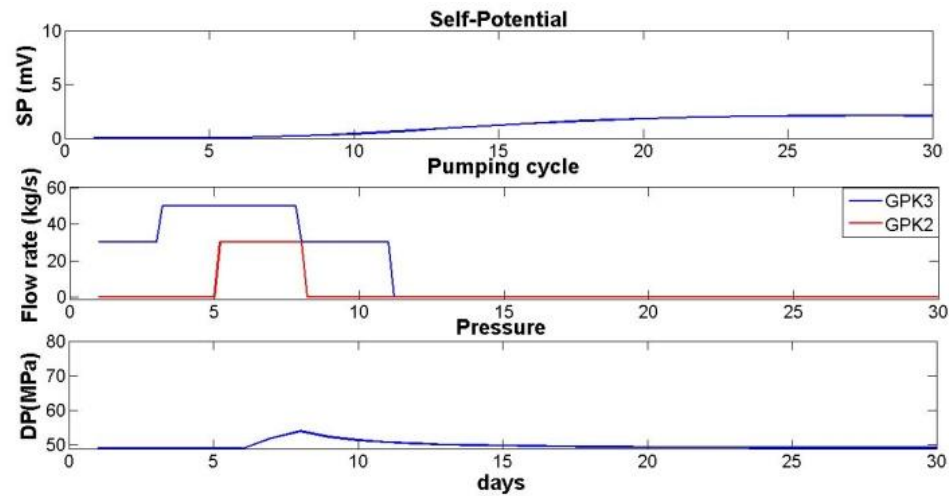
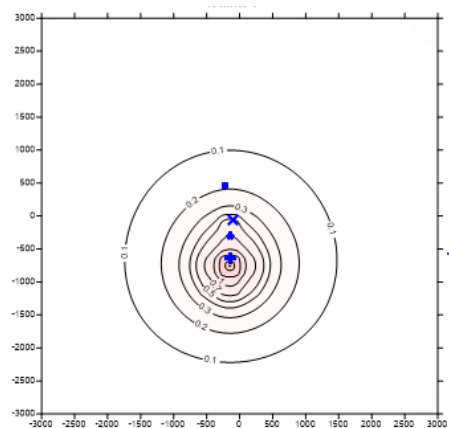
Second case: time dependent trend of electric potential



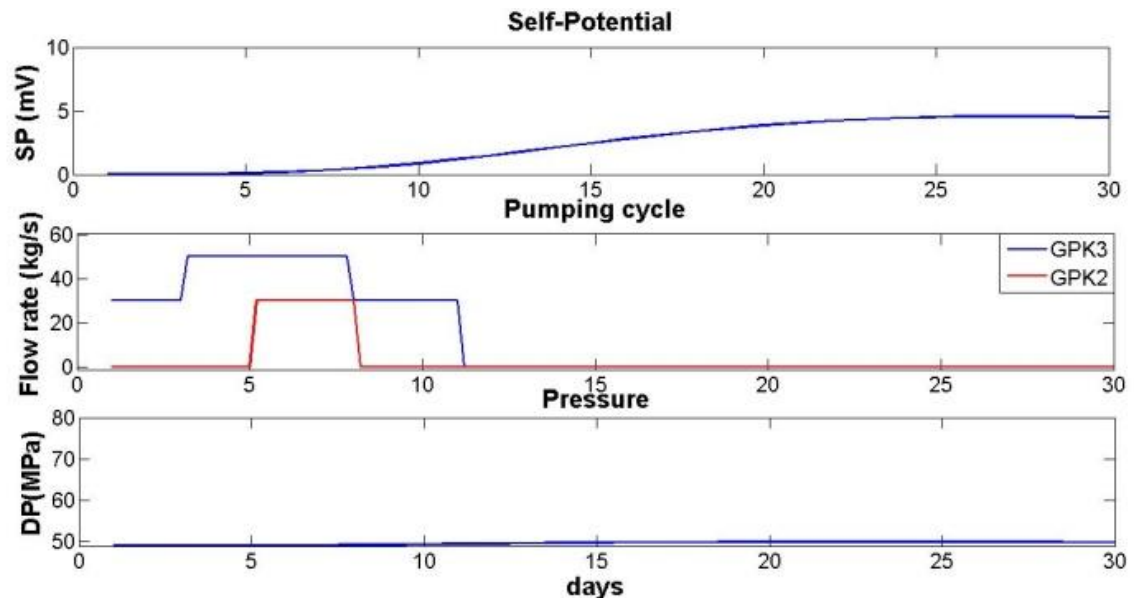
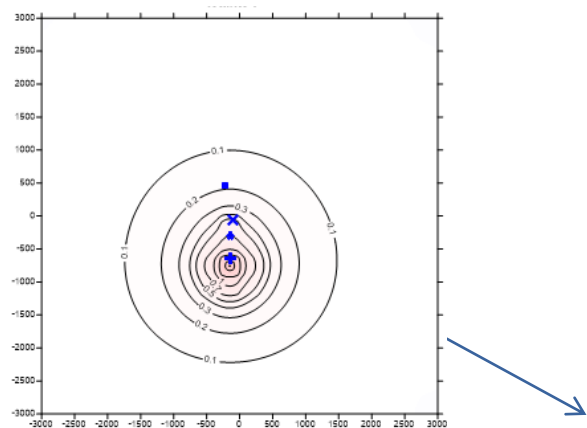
Second case: time dependent trend of electric potential



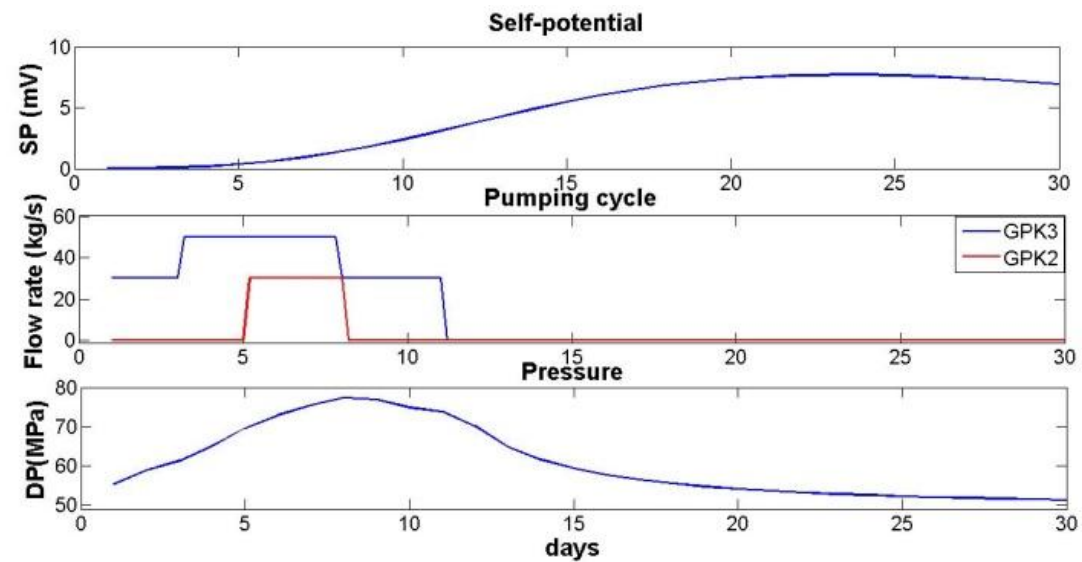
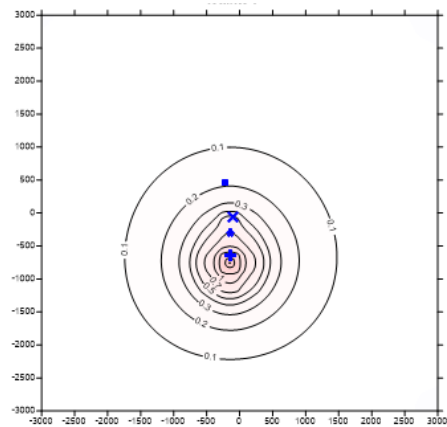
Second case: time dependent trend of electric potential



Second case: time dependent trend of electric potential



Second case: time dependent trend of electric potential



Conclusion

- ▶ Evaluation with the synthetic case of the model created in Comsol
- ▶ Pressure shows a rapid decreasing trend just after the closure of the wells

Conclusion

- ▶ Electric potential reaches the maximum slowly and it shows a relaxing phase many days after the closure of the wells
- ▶ Electric potential follows better than pressure the groundwater flow
- ▶ Self-Potential explains the occurrence of induced seismicity many days after the closure of the wells by the still presence of water

THANKS FOR ATTENTION